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CORRELATION OF ARIZONA PALEOZOIC FORMATIONS

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TUCSON SUB-AREA

...e, and Cochise formation.—At the base of the Tucson sub-area, is the Bolsa quartzite, a buffly gritty, and cross-bedded, vitreous quartzite, a total thickness varying from 430 feet to 500 feet. The Picacho de Calera Hills, 25 miles west of Tucson, is Pima sandstone, a buff, hard sandstone, and quartzite in places. This sandstone, only more than 60 miles, from the Whetstone Mountains to the Picacho de Calera Hills, Pima County, contains *Iphidella pannula* (White).

...and the Abrigo is the Cochise formation. In the Whetstone Mountains, it consists of three divisions. The upper division is pink and reddish, thin-bedded sandstone, which follows 116 feet of yellow, pink, gray, and buff, which in the upper horizons alternates with buffly limestone. Brachiopods, including *Neolenus intermedius pugio* Walcott, are abundant in some abundance, and *Neolenus* n. sp.) have been found in the upper division is characterized by blue limestone, rather grayish, brownish, and buff, and higher in the section, the blue limestone is more numerous. Within the lower part of the blue limestone, with them, are thin-bedded, calcareous sandstone.

...*Neolenus intermedius pugio* Walcott, and fragmental. Layers with edgewise bedding in the upper division is easily recognizable, since they weather more than the beds of the lower division. The formation are beds of oolitic limestone, but yet thick, replete with small spherules, and showing to show algal structure under a microscope. The limestone are thin-bedded pink and gray, and thin-bedded gray arenaceous limestone. Specimens of *Alokistocare* cfr. *linnaei* Walcott. These beds are of interest for the presence of *Alokistocare*, all undescribed specimens of a *Micromorphus* (P.) *superba* (Walcott) principally in the pygidium. The top of the formation contains the algae-like inclusions. This upper division is 165 feet in thickness, and the thickness of the entire Cochise formation in the Whetstone Mountains is 311 feet.

division is 165 feet in thickness, and the thickness of the entire Cochise formation in the Whetstone Mountains is 311 feet.

Abrigo formation.—In the type locality at Bisbee, the Abrigo has been characterized by Ransome as "thin-bedded impure, in part shaly, in part arenaceous, very cherty dolomitic limestone. Carries Middle Cambrian fossils. Bed of white quartzite about 8 feet at top."¹⁷ The thickness is given as 770 feet. The age was determined by Walcott, on the basis of the paleontologic material collected by Ransome. No species names, however, were given in any of Ransome's publications. Later, Walcott cited the following brachiopods from the Abrigo: *Obolus tetonensis* Walcott and *Obolus zetis* (Walcott).¹⁸ In a still later paper, a cranium and a pygidium of *Crepicephalus texanus* (Shumard), found in the Moore Canyon near Bisbee, are mentioned.¹⁹

Obviously, the part of the Cambrian limestone of Bisbee that contains *Obolus zetis* is older than the beds with *Crepicephalus texanus*. Although no Middle Cambrian trilobites have been found in the Bisbee area thus far, the "algal" limestone so characteristic of the Cochise formation is present in the Escabrosa Ridge west of Bisbee.

The base of the Abrigo formation is marked by the sudden appearance of *Crepicephalus texanus* (Shumard) and *Hesperaspis butleri* n. sp., the first-named form persisting throughout the entire formation, the second seemingly restricted to its basal layers. So characteristic is the presence of *Hesperaspis butleri* at the base of the Abrigo that in every Cambrian outcrop in southeastern Arizona known to the writer this species has been collected. It occurs invariably with *Crepicephalus texanus*. The base of the Abrigo has been located in the Escabrosa Ridge at Bisbee; in the Whetstone Mountains, Cochise County, between Benson and Tucson; in the Picacho de Calera Hills, Pima County, 25 miles west of Tucson; and in the Santa Catalina Mountains, Pinal County, 30 miles north of Tucson.

When first found, the specimens of *Hesperaspis* were supposed to belong to a new genus. Realizing that similar material might be present in the collections of Walcott, or even described in Walcott's unpublished manuscripts, the writer in 1928 sent photographs of the Arizona form to David White, and asked him to submit them to Ulrich and Resser, who had charge of Walcott's material. White's answer contained the following statement:

Doctor Ulrich has examined your photographs of the trilobite found with *Crepicephalus texanus*. He says the form you have is allied to *Giordanella* from Sardinia,

¹⁷ L. Ransome: *Some Paleozoic sections in Arizona and their correlation*, U. S. Geol. Surv., Bull. 98-K (1916) pl. 25.

¹⁸ D. Walcott: *Cambrian Brachiopoda*, U. S. Geol. Surv., Mon. 51 (1912) p. 417, 422.

¹⁹ D. Walcott: *Cambrian geology and paleontology: Cambrian trilobites*, Smithsonian Misc. Coll., vol. 3 (1916) p. 213.

but is more strictly congeneric and is probably specifically the same as that described in unpublished manuscripts of Mr. Walcott as *Hesperaspis slateri* Walcott. This species occurs toward the close of the Crepicephalus fauna in the Rocky Mountains of Montana and in Wyoming.

Later the writer had an opportunity to examine the material from the Little Rocky Mountains, preserved in the Smithsonian Institution. The Arizona material is much better preserved, and its stratigraphic position is precisely known. At present, two species are distinguished in the Upper Cambrian of Arizona at the base of the Crepicephalus zone. To avoid *nomina nuda*, these forms are here provisionally designated as *Hesperaspis ransomei* n. sp., and *H. butleri* n. sp., both being new fossils for the basal Abrigo.

Hesperaspis ransomei, new species

(Plate 1, figures 6, 7)

Cranidium from medium size to large, somewhat flat, with a broad base, subrectangular above the postero-lateral limbs, highest above its mid-length, the anterior part of the glabella and in front of the eyes, thence it slopes gently anteriorly and posteriorly. Glabella very faintly outlined but nearly always visible in reflected light or when covered with ammonia powder. The shallow and obtuse dorsal furrows are strongest in the posterior part of the glabella and almost invisible in front; consequently, the glabella is distinctly separated from the fixed cheeks in its posterior part only, where it is raised above the postero-lateral limbs and elevated along a gentle and obtuse median ridge which disappears anteriorly, whereas, anteriorly it merges into the frontal limb. Nevertheless, the rounded-in-front outline of the glabella is clearly indicated. There are no glabellar furrows. The broad, convex, crescentic, occipital segment, separated from the glabella by a very shallow, barely perceptible depression representing the occipital furrow, bears a strong median elevation, passing posteriorly into a sharp occipital spine.

Postero-lateral limbs large, wide, semicrescentic in outline, sloping outward and posteriorly, abutting against the posterior part of the glabella, with deeply impressed posterior furrows not connected with the occipital furrow. Fixed cheeks, less than half as wide as the glabella, merge into a broad frontal limb which is separated by an almost imperceptible depression from the narrow frontal border. Palpebral lobes of medium size, not elevated above the fixed cheeks, contained about two times in the length of the glabella, are placed about the mid-length of the cranidium. Palpebral ridges absent.

Facial sutures, cutting the posterior border within the genal angles, ascend gently around the postero-lateral limbs to the base of the palpebral lobes, outline the latter, cutting a little inward above, and thence extend anteriorly and outward, thus giving an expanding aspect to the frontal limb. They curve rather abruptly inward at the frontal border and are intramarginal for a short distance.

Free cheeks rather wide, slightly convex, with genal angles extended backward in strong short spines.

Thorax unknown.

Pygidium from semi-elliptical to subtrigonal in outline, short and broad. The convex subcylindrical axial lobe is elevated, especially so in its posterior part, and does not reach the posterior extremity. On the axial lobe, the three anterior axial annulations are very distinct, the first one more raised than the rest. The fourth annulation is distinct. On some specimens, a faint fifth annulation may be distinguished between the fourth annulation and the elevated terminal section. The axial furrows are well developed on the sides of the axial lobe, whereas posteriorly there is a slope inclined about 45 degrees from the top of the terminal section to the end of the posterior extremity, with the base of the terminal section scarcely separable from the posterior border. Only the first and the second outward expansions of the anchylosed segments of the axial lobe are indicated on the pleural lobes, by the narrow furrows of which the first is quite well marked. The pleural lobes and the posterior border are outlined by a narrow, flattened peripheral rim.

Hesperaspis butleri

(Plate 1, figure 8)

This species differs from *Hesperaspis ransomei* in its inflated cranidium, highest in the center, and the glabella always visible in front, and in the differentiation of the frontal border, which is closer to the glabella, a stronger occipital spine.

H. butleri is a smaller form and never attains the size of *H. ransomei*.

Hesperaspis butleri is widely distributed in practically every known outcrop of the Escabrosa Canyon between Bisbee and the Picacho de Calera Hills; and in the base of the Abrigo formation with

the Whetstone Mountains, appearing

in the Abrigo with *Hesperaspis butleri*

made up of trilobite fragments, and

in the Whetstone that have casts of large

trilobites here in association with a number

of other genera. *Asaphiscus (Blainia)* ;

forms related to *Asaphiscus*, but

which do not characterize this horizon.

Copper Queen limestone.—The yellow

limestone of northern Arizona consists of localities

in the Bisbee-Tucson sub-area, the

Escabrosa limestone formation is represented by

the Escabrosa limestone and the Rincon limestone

of Ransome²⁰ has described in the

Escabrosa limestone 6 feet in thickness above

the Escabrosa limestone withstands weathering

in the Abrigo, which is made of "rather

soft limestone." In Ransome's Figure 4, r

is shown in the Escabrosa Ridge n

exactly as the upper hill within t

of trilobites collected there by J

Maladia Walcott and *Irvingella* ?

A close examination of this limestone

is here called the Copper Queen

mining area of Bisbee, w

defined, this formation embraces

²⁰ F. L. Ransome: *Description of the Bisbee*

Folio no. 112 (1904) p. 3.

See also N. H. Darton: *A résumé of*

Arizona p. 47.

N. H. Darton: *op. cit.*, p. 50.

Hesperaspis butleri, new species

(Plate 1, figure 8)

This species differs from *Hesperaspis ransomei* in having a much-narrower and more-inflated cranium, highest in the middle, with a much-longer and better-outlined glabella always visible in front, a shorter and narrower frontal limb without any differentiation of the frontal border, narrow fixed cheeks bringing the palpebral lobes closer to the glabella, a stronger-developed occipital furrow, and a wider occipital spine.

This is a smaller form and never attains the size of larger specimens of *Hesperaspis ransomei*.

Hesperaspis butleri is widely distributed in southeastern Arizona. It has been found in practically every known outcrop of the Cambrian: at Bisbee, in the Tombstone Canyon between Bisbee and Tombstone; in the Whetstone Mountains; in the Picacho de Calera Hills; and in the Santa Catalina Mountains. It is found at the base of the Abrigo formation with *Crepicephalus texanus* (Shumard).

In the Whetstone Mountains, approximately 240 feet above the base of the Abrigo with *Hesperaspis butleri*, occurs a gray limestone practically made up of trilobite fragments, and containing angular inclusions of red sandstone that have casts of large *Hyolithes*. *Crepicephalus texanus* is found here in association with a number of new species and varieties of this genus. *Asaphiscus (Blainia) gregarius* Walcott and a large variety of forms related to *Asaphiscus*, belonging mostly to *Blountia* Walcott, characterize this horizon.

Copper Queen limestone.—The youngest part of the Cambrian in southeastern Arizona consists of localized facies, each very characteristic. In the Bisbee-Tucson sub-area, the Upper Cambrian younger than the Abrigo formation is represented by two different facies, the Copper Queen limestone and the Rincon limestone.

Ransome²⁰ has described in the Bisbee area "one bed of harder gray limestone 6 feet in thickness about 40 feet from the top" of his Abrigo limestone withstands weathering much better than the rest of the Abrigo, which is made of "rather soft, sandy, thin-bedded, gray limestone."

In Ransome's Figure 4, representing the Mount Martin Paleozoic section in the Escabrosa Ridge near Bisbee, this limestone is shown distinctly as the upper hill within the Abrigo limestone. A few specimens of trilobites collected there by Ransome were referred by Walcott²¹ to *Blountia* Walcott and *Irvingella* Ulrich and Resser.

A close examination of this limestone revealed its stratigraphic importance. It is here called the Copper Queen limestone, from the Copper Queen mining area of Bisbee, where it is a good key horizon. As here defined, this formation embraces much more than 6 feet of hard limestone.

²⁰Ransome: *Description of the Bisbee quadrangle, Arizona*, U. S. Geol. Surv., Geol. Atlas, sheet no. 112 (1904) p. 3.

²¹See N. H. Darton: *A résumé of Arizona geology*, Univ. Ariz., Bull. 119, Geol. Ser. no. 3.

Darton: *op. cit.*, p. 50.

EXPLANATION OF PLATE

PLATE 1

DEVONIAN FISHES AND UPPER CAMBRIAN TRILOBITES

FIGURES 1-5.—Plates of primitive *Arthrodira* from Devonian beds of Mt. Elden near Flagstaff, Arizona. (1) Dorsomedian plate, convex, seen from the outside. Superficial layers are partially denuded and inferior keel is slightly exposed. $\times 2$. (2)—Posterior ventrolateral plate. Cast. $\times 2$. (3)—Anterior ventrolateral plate? Cast. Lateral lines are seen. Natural size. (4)—Dorsomedian plate, concave, seen from the inside. It shows the inferior keel. $\times \frac{1}{2}$. (5)—Posterior ventrolateral plate. Cast. Natural size.

FIGURE 6.—*Hesperaspis ransomei* sp. n. Holotype, cranidium.

FIGURE 7.—*Hesperaspis ransomei* sp. n. Holotype, pygidium.

FIGURE 8.—*Hesperaspis butleri* sp. n. Holotype, cranidium.



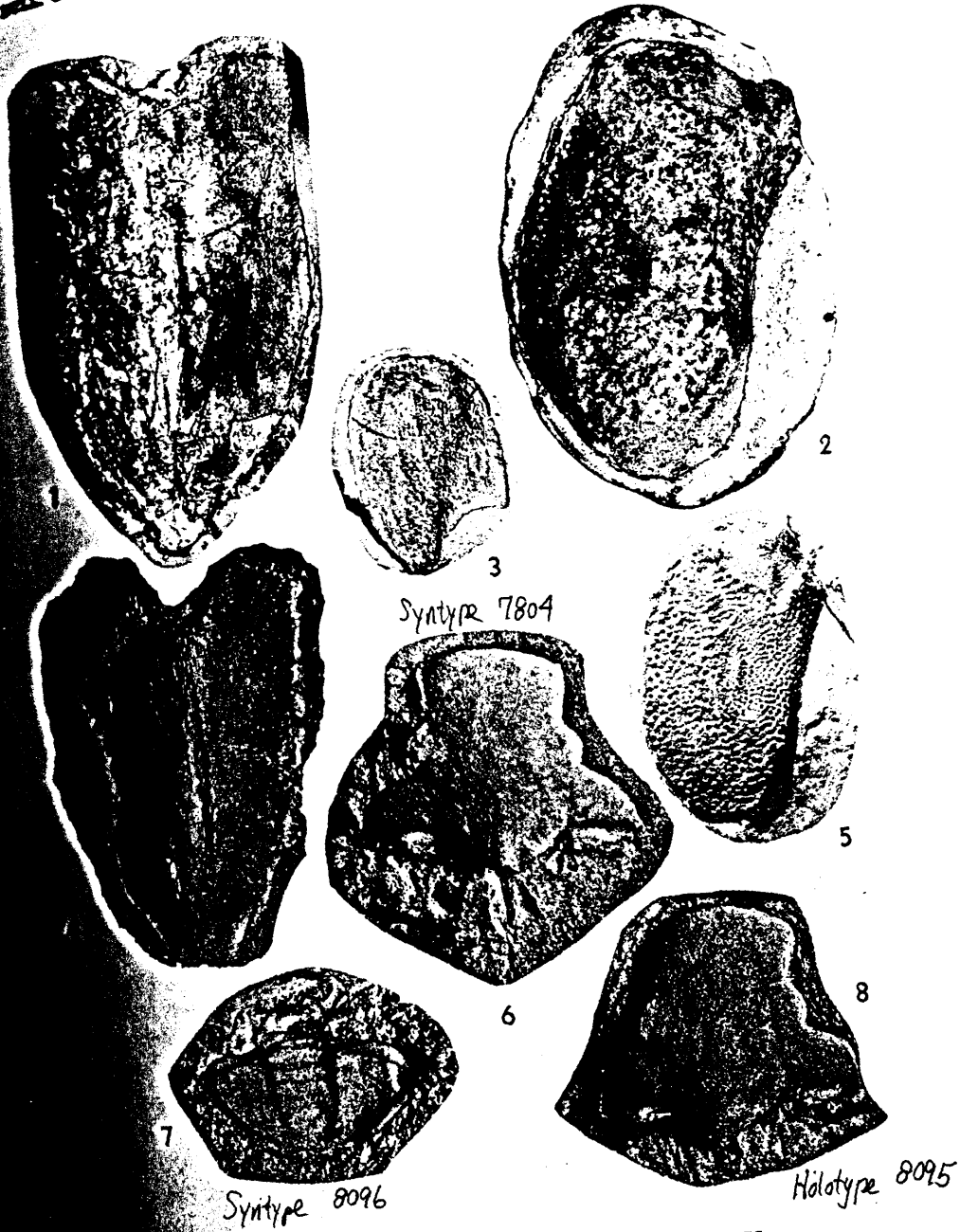
DEVONIAN FISHES AND UPPER CAMBRIAN TRILOBITES
 Figures 1-5. Plates of primitive *Arthrodira* from I
 Figures 6, 7. *Hesperaspis ransomei* sp.
 Figure 8. *Hesperaspis butleri*

PLATE

CAMBRIAN TRILOBITES

from Devonian beds of Mt. Elden.
 median plate, convex, seen from the
 tially denuded and inferior keel
 anterior ventrolateral plate. Cast. X
 Cast. Lateral lines are seen
 ate, concave, seen from the inside
 (5)—Posterior ventrolateral plate

cranidium.
 pygidium.
 cranidium.



DEVONIAN FISHES AND UPPER CAMBRIAN TRILOBITES

1-5. Plates of primitive *Arthrodira* from Devonian beds of Mt. Elden, near Flagstaff, Arizona.
 Figures 6, 7. *Hesperaspis ransomei* sp. n. holotype, (6) cranidium (7) pygidium.
 Figure 8. *Hesperaspis bulleri* sp. n., holotype, cranidium.