

Figure 16. Close-ups of outcrops of outer shelf to slope channel facies shown in Figure 15. A) Base of lower amalgamated sandstone with a few scattered ripup clasts; scale is 1.5 m. B) Base of upper medium sandstone body grading into laminated fine sandstone with flame structures. Above the flame structures, the fine sandstone grades into laminated very fine sandstone containing small channels filled with coarser material; coin is 2.4 cm in diameter.

PALEOENVIRONMENTAL SUMMARY

The late early through early middle Eocene Llajas Formation of the Simi Valley area represents a transitional alluvial to marine system. In a vertical sequence, most of the formation is transgressive (retrogradational) in that shallower deposits are overlain by deeper-water deposits. Based on regional paleogeographic analyses (Howell,

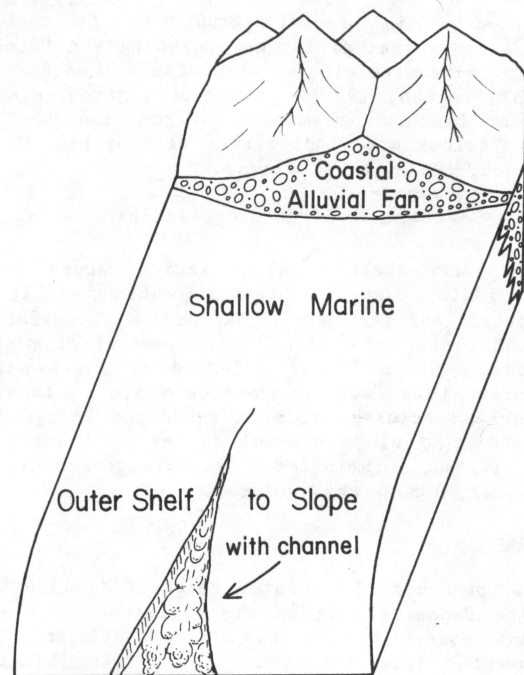


Figure 17. Depositional model of the Eocene Llajas Formation.

1975; Nilsen and McKee, 1979), the shoreline was probably oriented in a north-south direction with a highland to the east. Topography most likely resembled modern high-relief areas along the west coast of Baja, California. Gravelly and sandy alluvial-fan deposits accumulated along the shoreline (Fig. 17). Source rocks were probably nearby Paleocene deposits. During storms, braided rivers of the alluvial-fan system carried coarse sediment into the transgressing sea. Shoreline deposits consisted of gravel interfingering with well laminated, well sorted, *Ophiomorpha*-burrowed sands containing abundant, mostly fragmented shells of shallow-marine mollusks.

Just offshore from these rivers, sedimentation was rapid in the shallow-marine environment, and sand build up caused shoaling conditions. The shallow-marine (transgressive) deposits were predominantly laminated sands alternating with bioturbated sands. The laminated sands represent storm-influenced stratification, and the bioturbated sands represent inter-storm activity of burrowing organisms. Well preserved indigenous shallow-marine mollusks and benthic foraminiferal assemblages were present but were largely confined to surge-channel deposits. In more seaward areas, the deeper the water the less amount of laminated sands and the more amount of bioturbated sands. Where such shallow-marine deposits graded into the outer shelf to slope environment, sediments consisted mostly of bioturbated silty sands and the invertebrate remains show less signs of reworking. One layer, the "Stewart bed", contains an essentially in situ megafaunal assemblage that consists predominantly of carnivore and herbivore gastropods and shallow-burrowing bivalves. Extant genera of the "Stewart bed", as well as the previously mentioned indigenous molluscan assemblages, most commonly occur today in warm temperate to tropical seas less than 80 m depth.

Deposits which formed in the outer shelf to slope area consisted of bioturbated siltstone.

Molluscan remains were scarce, but benthic foraminiferal assemblages are diverse and indicative of the slope environment. Locally, these deposits were incised by turbidity currents. The resulting small-scale channels were subsequently filled with amalgamated turbidite sandstone with minor amounts of gravel containing displaced shallow-marine mollusks.

The uppermost part of the formation represents a regressive (progradational) event in which the outer shelf to slope channel-fill deposits and enclosing outer shelf to slope deposits were covered by shallower deposits. Where such outer shelf to slope deposits graded into these shallower deposits, sediments consisted of bioturbated silts and silty sands. Shallow-marine mollusk and benthic foraminiferal taxa are similar to those found where the transgressive shallow-marine deposits grade into the outer shelf to slope deposits.

The regressive cycle is incomplete due to truncation by erosion prior to the deposition of the overlying nonmarine early late Eocene through Oligocene Sespe Formation.

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