



Fig. 4.—*Icriocarcinus xestos* Bishop, 1988. A, transverse and B, oblique views of paratype, SDSNH 26040, preserved within a complex, sand-filled burrow. The position of the specimen is shown by the claws and the sternum (S), seen in the oblique view (B). The lower continuation of the burrow (B) is also visible on the oblique view.

cle completely exfoliated, so that the dorsal carapace surface appears smooth and the grooves appear very deep. The other style is with preserved cuticle, which exhibits small granules that can become very closely spaced to appear like wrinkles or scabrous ridges. In this style of preservation, the grooves are still apparent but are not as deep as in specimens lacking cuticle.

As discussed above, *Icriocarcinus* is morphologically quite similar to *Ommatocarcinus*. It is also noteworthy that species of both genera are burrowers that produce complex burrow structures. Both fossil and extant members of *Ommatocarcinus* are known as burrowers

(Jenkins 1975). Jenkins (1975) described exposed burrows, containing well-preserved remains of *Ommatocarcinus corioensis* (Cresswell, 1886) from the lower Miocene-Pliocene Port Campbell Limestone, southeastern Australia and Tasmania. The burrows are exposed on bedding planes, and single burrows may extend for a distance of over a meter. As described and illustrated (Jenkins 1975, pl. 8), the burrows are branching structures in which horizontal elements dominate over vertical ones. The exposed surfaces of the burrows seem to be either smooth or pelleted.

One paratype of *Icriocarcinus* recovered from the

Point Loma Formation was found preserved in a well-defined burrow structure (Fig. 4). That burrow, originally illustrated by Bishop (1988, fig. 3D), appears to be relatively short and contains a complete specimen of *Icriocarcinus xestos* within it. The burrow is constructed in a grey mudstone containing sand-sized particles of shell fragments and organic debris. The burrow is filled with fine sandy sediment with large fragments of mollusks. If the burrow is oriented with the preserved crab in a horizontal position, the uppermost part of the burrow descends into the sediment at a 60° angle to a depth of about 3.5 cm, at which point the burrow continues into the sediment at a 35° angle to a depth of about 8 cm. The structure is broader in the upper segment with a maximum measurable dimension of 4 cm, whereas the lower segment has a maximum measurable dimension of about 2.5 cm. These measurements provide relative sizes but the cross-section of the burrow and the maximum diameter cannot be determined because the structure is viewed along the broken section. The crab appears to be positioned within a chamber more than 10 cm wide and 4.5 cm high. The sediment filling the chamber is darker in color but otherwise similar to the mudstone in which the burrow is constructed.

Examination of the lower surface of the specimen, however, indicates that the structure is more complex. On this surface, the sternum of the animal is exposed along with a continuation of the sand-filled burrow structure (Fig. 4). The downward extension of the burrow structure is ovoid with a maximum diameter of 5 cm and a minimum diameter of 3.5 cm. Thus, as in the case of the burrows of *Ommatocarcinus*, the burrow of *Icriocarcinus xestos* is complex.

Ichnofossils in the Point Loma Formation have been studied previously (Kern and Warme 1974). Burrow structures were recognized not unlike the one described above in addition to two observations that are relevant to this study. Kern and Warme noted (1974, p. 896, fig. 6) that when the burrows passed from sandy sediment into muddy sediment, the burrow angle decreased from being vertically dominated to becoming horizontally dominated. Further, they noted that the burrows were pellet-lined in the sandy sediment, and would likely be referred to *Ophiomorpha*; whereas the same burrow in muddy sediment was smooth and unlined, so that it would be referred to the ichnogenus *Thalassinoides*. The burrow in which *Icriocarcinus xestos* was preserved is steep at the top, becoming more gently sloping downward and is sand-filled. Perhaps the structure originated in sandy sediment, ultimately penetrating muddy sediment.

The environment of deposition of the Point Loma Formation has been interpreted to be at bathyal depths at which mudstones were deposited by relatively slow processes of sedimentation whereas the sandstones were accumulated rapidly by grain-flow processes (Kern and Warme 1974). Interpretation of water depth as well as

the downslope accumulation of sediments was based upon the mixing of sublittoral and bathyal foraminiferans (Sliter 1968 in Kern and Warme 1974). Thus, *Icriocarcinus* may have inhabited water deeper than, or as deep as, that in which *Ommatocarcinus* has been reported. Sakai (1976) recorded *O. macgillivrayi* White, 1851, from depths up to 100 m, and Jenkins (1975) reported the same species at depths of 274 m. Another species, *O. fibriophthalmus* Yokoya, 1933, was collected at a depth of 146 m (Sakai 1976).

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